Development of $q_c$ – $V_s$ Correlation for Depok Silt-Clay

Widjojo A. Prakoso
Civil Engineering Department University of Indonesia, Depok 16424, E-mail: wprakoso@eng.ui.ac.id

Abstract

To initiate the development of cone penetration resistance $q_c$ and shear wave velocity $V_s$ correlations of soils from Indonesia, the results of mechanical cone penetration soundings and seismic downhole tests from Depok, West Java area are compiled and analyzed. The soils are predominantly silt-clay soils. The results of the tests are used to develop a $q_c$ – $V_s$ correlation. The results and the associated correlation are then compared to $q_c$ – $V_s$ correlations from other countries.

Keywords: Shear wave velocity, cone penetration test, cone resistance.

1. Introduction

The shear wave velocity of soils plays an important role in the design of geotechnical structures under dynamic loads. It is used mostly for determining the seismic site categories (e.g., BSN, 2002) and for an initial reference value for large strain problems related to seismic loading. In Indonesia, the shear wave velocity is typically measured using the seismic downhole test. However, the equipment is not widely available and, consequently, the test is generally too expensive to perform for most construction projects. On the other hand, mechanical cone penetration tests are the most common in-situ test because it is light-weight and easy to perform.

No direct cone penetration resistance $q_c$ and shear wave velocity $V_s$ correlations of soils from Indonesia are currently available. Furthermore, unlike correlations for standard penetration test results N-SPT and $V_s$, the available direct $q_c$ – $V_s$ correlations from other countries are limited (Andrus et al., 2003; Madiai and Simone, 2004; and Sun et al., 2008). Therefore, it is of interest to develop a $q_c$ – $V_s$ correlation of soils from Indonesia. To initiate this development, $q_c$ and $V_s$ data from the University of Indonesia complex in Depok, West Java are analyzed. In this paper, the database of the two parameters and the analysis conducted are described, followed by a discussion on the proposed correlation and on the comparison of $q_c$ – $V_s$ correlations.

2. Test Program

The test program considered in this study consisted of three locations (A, B, and C), with one mechanical cone penetration test (CPT) and one seismic downhole test (SDHT) conducted at each location. The locations were within the University of Indonesia complex in Depok, West Java. The distance between locations A and C was about 250 m, while the distance between locations B and C was about 180 m.

The mechanical CPTs were conducted in accordance with ASTM D3441 (2008). The cone with an apex angle of 60° is 10 cm$^2$ in cross-sectional area and has a 150 cm$^2$ friction sleeve. The cone penetration resistance $q_c$ and the associated friction ratio $R_f$ readings and calculations were taken and performed at 0.2 m interval. The Robertson (1990) procedure was modified to further interpret the CPT results; the normalized cone resistance $Q$ and the normalized friction ratio $F$ respectively are given by the following:

$$Q = \frac{(q_c - s_v)}{s'_v}$$  (1)
$$F = \frac{f_r}{(q_c - s_v)}$$  (2)
Development of $q_c - V_s$ Correlation for Depok Silt-Clay

in which $\sigma_v = \text{total vertical stress}$ and $\sigma'_v = \text{effective vertical stress}$, and $f_s = \text{sleeve friction}$.

The seismic downhole tests were conducted using OYO Borehole Pick Model 3315 and McSeis-SX 48 Model 1126C. The shear wave velocity was measured at 1.0 m interval.

3. Test Results

3.1 Location A

The cone penetration resistance $q_c$ and the friction ratio $R_f$ of the cone penetration test (CPT) for Location A are shown in Figure 1. The normalized cone resistance $Q$ and the normalized friction ratio $F$ are also shown in Figure 1. The shear wave velocity $V_s$ from the seismic down-hole test (SDHT) in an adjacent deep boring is shown in Figure 1 as well. Based on the results, three geomaterial layers can be identified: ① depth = 0 – 3.0 m, ② depth = 3.0 – 7.0 m, ③ depth = 7.0 – 18.0 m, and ④ depth = >18.0 m. In addition, the groundwater table in the deep boring was found at a depth of 6.2 m.

The Robertson (1990) procedure was used to further interpret the CPT results of the four layers. As shown in Figure 2, the first layer is predominantly in Zone 3 (clay to silty clay) with higher overconsolidation ratio (OCR), the second layer is predominantly in Zone 4 (clayey silt to silty clay), the third layer is predominantly in Zone 3 with lower OCR, and the fourth layer is predominantly in the Zone 4 and Zone 5 (silty sand to sandy silt) with relatively low OCR. It is noted that the particle size analysis (ASTM, 2002) of an undisturbed sample from depth of 5.5 – 6.0 m resulted in clay = 24%, silt = 75%, sand = 1%, while that of an undisturbed sample from depth of 14.0 – 14.5 m resulted in clay = 23%, silt = 76%, sand = 1%.

Figure 1. Mechanical CPT and SDHT results for Location A

Figure 2. Q-F analysis for Location A

(■: 0-3.0 m, ×: 3.0-7.0m, ◊: 7.0-18.0m, ◊: >18.0m)
3.2 Location B

The $q_c$ and $R_f$ profiles of the CPT for Location B are shown in Figure 3. The Q and F profiles are also shown in Figure 3. The $V_s$ from the SDHT in the adjacent deep boring is shown in Figure 3 as well. Based on the results, three geomaterial layers can be identified: ① depth = 0 – 3.0 m, ② depth = 3.0 – 7.5 m, ③ depth = 7.5 – 20.0 m, and ④ depth >20.0 m. In addition, the groundwater table was found at a depth of 9.1 m.

Based on the Robertson (1990) procedure shown as Figure 4, the first through the third layers are predominantly in Zone 3 (silty clay to clay) with decreasing OCR, and the fourth layer is predominantly in the Zone 4 (clayey silt to silty clay) and Zone 5 (silty sand to sandy silt) with relatively low OCR. It is noted that the particle size analysis (ASTM 2002) of an undisturbed sample from depth of 1.5 – 2.0 m resulted in clay = 23%, silt = 75%, sand = 2%, that of an undisturbed sample from depth of 8.5 – 9.0 m resulted in clay = 25%, silt = 74%, sand = 1%, and that of an undisturbed sample from depth of 17.0 – 17.5 m resulted in clay = 23%, silt = 76%, sand = 1%.

Figure 3. Mechanical CPT and SDHT results for Location B

Figure 4. Q-F analysis for Location B
(■: 0-3 m, ×: 3.0-7.5m, ◇: 7.5-20.0m, ◇: >20m)
3.3 Location C

The $q_c$ and $R_f$ profiles of the CPT for Location C are shown in Figure 5. The $Q$ and $F$ profiles are also shown in Figure 5. The $V_s$ from the SDHT in the adjacent deep boring is shown in Figure 5 as well. Based on the results, three geomaterial layers can be identified: ① depth = 0 – 3.0 m, ② depth = 3.0 – 19.0 m, and ③ depth >19.0 m. In addition, the groundwater table was found at a depth of 6.2 m.

Based on the Robertson (1990) procedure shown as Figure 6, the first and the second layers are predominantly in Zone 3 (silty clay to clay) with decreasing OCR, and the third layer varies between Zone 3 and Zone 5 (silty sand to sandy silt). It is noted that the particle size analysis (ASTM 2002) of an undisturbed sample from depth of 3.0 – 3.5 m resulted in clay = 23%, silt = 75%, sand = 2%, while that of an undisturbed sample from depth of 10.0 – 10.5 m resulted in clay = 23%, silt = 73%, sand = 4%.

4. Correlation Development and Discussion

A simple regression analysis was performed, taking the shear wave velocity $V_s$ as the dependent parameter and the cone penetration resistance $q_c$ as the independent parameter. As the $q_c$ and $V_s$ were determined at different intervals, five $q_c$ values had to be averaged for the associated depth of $V_s$ value. It is noted that the upper 3.0 m of $V_s$ values at Location A, the upper 1.0 m of $V_s$ values of Location B, and the upper 2.0 m of $V_s$ values of Location C were not included in the analysis, as they appeared to be unusually high for relatively low $q_c$ values.

Figure 7a compares $q_c$ and $V_s$ data for different materials based on the Robertson (1990) criteria. It can be observed that all data cluster in the same general range, and it can be concluded therefore that material types would not have a significant effect on the correlation.
Figure 7a also shows the initial $q_c - V_s$ correlation equation based on 56 data points. The distribution of the $V_s$ deviation from this initial correlation equation is shown as Figure 7b, and the standard error is 27.39 m/s. There are four data points that deviate significantly (> 60 m/s), and these data points are considered as outliers in the further analysis.

Based on the select $V_s$ and $q_c$ data, the following correlation has been derived for the Depok silt-clay:

$$V_s = 115.70 (q_c)^{0.34} \quad (3)$$

in which $V_s$ in m/s and $q_c$ in MPa, and the standard error is 18.16 m/s. Note that the number of data is 52 and the $r^2$ value is 0.69. The data and the regression line and equation are shown in Figure 9, while the distribution of the $V_s$ deviation from the proposed correlation equation is shown as Figure 10.

The correlation above is compared with similar correlations from other countries. It is noted that direct $q_c - V_s$ correlations for clayey materials are limited. Andrus et al. (2003) proposed the following correlation for clayey holocene soils from the USA (South Carolina and California), Canada, and Japan:

$$V_s = 6.21 (q_c)^{0.444} \quad (4)$$

in which $q_c$ in kPa ($n = 31$, $r^2 = 0.83$). Madiai and Simone (2004) proposed the following correlation for some clayey soils from Italy:

$$V_s = 211.2 (q_c)^{0.231} \quad (5)$$

in which $q_c$ in MPa ($n = 46$, $r^2 = 0.871$). Sun et al. (2008) proposed the following correlation for some clayey soils from South Korea:

$$V_s = 17.84 (q_c)^{0.301} \quad (6)$$

in which $q_c$ in kPa ($r^2 = 0.741$). In addition, Mayne and Rix (in Mayne, 2007) developed a clayey soil database for the $q_c - V_s$ correlation. Although their database is a $q_c$ database, but this database can be used to compare the above-groundwater-table data.

The comparison is shown as Figure 11. It can be observed that the present $q_c - V_s$ data and the associated correlation equation are slightly lower than the correlation proposed by Andrus et al. (2003) for $q_c < 2$ MPa and are slightly lower than the correlation proposed by Sun et al. (2008) for $q_c > 2$ MPa. However, the present data and correlation equation appear to be in the lower bound of the Mayne and Rix’s database (in Mayne, 2007) and are significantly lower than the correlation proposed by Madiai and Simone (2004).
5. Conclusions

1. Cone penetration resistance \( q_c \) data obtained from mechanical CPTs and shear wave velocity \( V_s \) data obtained from seismic downhole tests from three locations within the University of Indonesia complex in Depok were evaluated.

2. The materials were predominantly silt-clay, determined using the Robertson’s criteria and confirmed by particle size analysis results.

3. Based on the results, a site-specific trend that correlates measured \( q_c \) to \( V_s \) for Depok silt-clay materials was developed.

4. The proposed correlation between \( V_s \) and \( q_c \) can be used for rough estimates of \( V_s \) from \( q_c \), particularly for preliminary studies and/or noncritical projects are under consideration.

5. From the comparison to similar correlations from other countries, the proposed correlation appears to be lower bound.

References


